

# Modeling the Albedo Advantage in Global Warming And an Albedo-Planck Parameter

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*(Please feel free to provide any helpful preprint comments to dfrsoft@gmail.com)*

**Key Words:** Re-Radiation Model, Global Warming Modeling, Planck Parameter, Planck-Albedo Parameter

**Abstract** In this paper, we model global warming using a re-radiation factor and the Planck's parameter to verify consistency. The re-radiation factor is important in quantifying the relative global warming impact of the albedo effect compared to the greenhouse gas (GHG) effects. The forcing due to the change in the Earth's global albedo compared to GHGs is found to have a 2.6 times larger impact on global warming. In our simple model, we additionally define a handy Planck-Albedo feedback parameter having a convenient value of  $1\text{W}/\text{m}^2/^\circ\text{K}/\Delta\%\text{albedo}$ . Using these results, it is concluded that a 1.5% solar geoengineering change in the global albedo could result in  $-4.8\text{W}/\text{m}^2$  of forcing. An alternate way to assess the Planck parameter was also found.

## 1 Introduction

Although global warming is highly complex, often it is helpful to work with a simplified model. We create a model that uses a re-radiation factor which helps to quantify significant differences between changes in the global albedo versus greenhouse gas forcing. We use the Planck's feedback parameter to verify model consistency. This model illustrates a reasonable way to view the Earth's energy budget; it provides a number of useful insights in climatology sensitivity estimates and demonstrates the relative advantage of solar geoengineering solutions over GHG reduction in global warming mitigation [1]. In working the model, we also find a handy Planck-Albedo parameter that may be useful to climatologists [2].

## 2. Data and Method

In order to introduce the re-radiation surface model, it is helpful to initially look at the Planck parameter as it plays a key role in verifying modeling.

### 2.1 Overview of Planck Feedback Parameter

Estimates on Planck's feedback parameter are varied, typically between  $-3.8\text{W}/\text{m}^2/^\circ\text{K}$  and  $-3.21\text{W}/\text{m}^2/^\circ\text{K}$  with some values as large as  $-7.1\text{W}/\text{m}^2/^\circ\text{K}$  [3]. The IPCC AR4 [4] list a value of  $-3.21\text{W}/\text{m}^2/^\circ\text{K}$ . Numerous authors have developed different expressions [3]. A typical estimate starts with

$$F_{\text{TOA}} = (1 - \alpha) S_o / 4 - \sigma (\beta T_s)^4 = (1 - \alpha) S_o / 4 - R_{\text{LWR}} \quad (1)$$

where  $S_o = 1361\text{W}/\text{m}^2$ ,  $F_{\text{TOA}}$  is the radiation budget at the top of the atmosphere,  $R_{\text{LWR}}$  is the outgoing long wave radiation (a function of surface temperature and albedo),  $\sigma$  is the Stefan-Boltzmann constant and  $\beta$  is described below. Then the Planck parameter  $\lambda_o$  can be calculated as

$$\lambda_o = \partial F_{\text{TOA}} / \partial T_s = -\partial R_{\text{LWR}} / \partial T_s \quad (2)$$

This result is

$$\lambda_o = -4\beta^4 \sigma T_s^3 = -4\beta \sigma T_{\text{toa}}^3 \quad (3)$$

where  $\beta$  varies from 0.876 to 0.887 (averaging=0.8815) and  $T_s = 288^\circ\text{K}$  [4]. This yields  $-3.37\text{W}/\text{m}^2/^\circ\text{K} < \lambda_o < -3.21\text{W}/\text{m}^2/^\circ\text{K}$ . However, from Eq. 3,  $\beta$  is often taken as the ratio

$$\beta = T_{\text{toa}} / T_s = 255^\circ\text{K} / 288^\circ\text{K} = 0.8854 \quad (4)$$

A common assessment uses  $T_{\text{toa}} = 255^\circ\text{K}$ , so that  $\lambda_o = -3.33\text{W}/\text{m}^2/^\circ\text{K}$ . Another expression developed by Schlesinger [5] is dependent on the albedo and surface temperature as

$$\lambda_o = S_o (1 - \alpha) / T_s \quad (5)$$

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When  $S_0=1361$ ,  $0.294118<\alpha<0.3$ , and  $T_s=288^\circ\text{K}$  then  $-3.308\text{W/m}^2/^\circ\text{K} >\lambda_o> -3.3358\text{W/m}^2/^\circ\text{K}$ , respectively.

## 2.2 Estimating Planck's Parameter with an Albedo Method

Consider a global albedo change corresponding to  $1^\circ\text{K}$  rise from solar absorption. Since we are only concerned with an albedo change

$$F_{TOA} = 0 = (1 - \alpha)E_o - \sigma(T_s)^4 \quad (6)$$

where  $E_o=S_0/4$ . Then a  $1^\circ\text{K}$  change is

$$\Delta T_s = T_2 - T_1 = \left( \frac{E_o}{\sigma} (1 - \alpha_2) \right)^{1/4} - \left( \frac{E_o}{\sigma} (1 - \alpha_1) \right)^{1/4} = 1^\circ\text{K} \quad (7)$$

Here we will use the AR5 albedo starting value of 0.294118 [6]. We find that the corresponding albedo change is 0.28299 when  $E_o=340\text{W/m}^2$ . This corresponds to an absorption of

$$\Delta E_o = E_o \{ (1 - \alpha_2) - (1 - \alpha_1) \} = E_o (\alpha_1 - \alpha_2) = 3.784\text{W/m}^2 \quad (8)$$

Since this is for a  $1^\circ\text{K}$  rise, then it can also be written as

$$\lambda_{1K} = 3.784\text{W/m}^2/^\circ\text{K} \quad (9)$$

We note this is related to the surface value, then

$$\lambda_{1K} = -4\sigma T_s^3 \quad (10)$$

By comparison to above we have

$$\lambda_o = \lambda_{1K} \beta = -3.784\text{W/m}^2/^\circ\text{K} = -3.349\text{W/m}^2/^\circ\text{K} \quad (11)$$

This is very close to the  $-3.33\text{W/m}^2/^\circ\text{K}$  value obtained in the traditional manner.

## 2.3 Top of the Atmosphere and Beta

From Eq. 1

$$R_{LWR} = \sigma(\beta T_s)^4 = \sigma(T_s)^4 \quad (13)$$

giving

$$\beta^4 R_{TOA, T_s} = R_{TOA, T_{TOA}} \quad (14)$$

We will need this expression later when showing model consistency with the Planck feedback parameter.

## 2.4 Re-radiation GHG GW Model

Global warming can be modeled by looking at two different time periods. We can model the radiation for 1950 as due to blackbody radiation with the addition of GHG re-radiation so

$$P_{Total, 1950} = P_\alpha + P_{GHG} = P_\alpha + f_1 P_\alpha \quad (15)$$

where  $P_\alpha = S_0 \{0.25x(1 - Albedo)\}$  and  $S_0=1361\text{W/m}^2$ . Here we have a fraction of the blackbody radiation is re-radiated by the GHGs so  $f_1$  is a re-radiation parameter. That is, the energy,  $P_{GHG}$ , must be some fraction  $P_\alpha$ . In 2019 due to global warming, this model is more complex and harder to separate out terms. However, it can still be done in theory, so

$$P_{Total 2019} = P_{\alpha'} + P_{GHG'+Feedback} = P_{\alpha'} + f_2 P_{\alpha'} \quad (16)$$

Here  $P_{GHG'+Feedback}$  includes GHG and its increase comprising also of water-vapor increase, lapse rate effect and other effects such as an increase in snow-ice albedo changes that are hard to separate out. That is, some of this feedback is related to GHG increases and some is related to albedo change.  $P_{\alpha'}$  represents any albedo change due to UHI absorption increases, cloud absorption change, ice and snow melting and so forth that can be discerned.

120 The re-radiation still must connect the absorption to re-radiation. We use a linear  $f$  parameter that indicates the  
 121 fraction of  $P_\alpha$  power that must be re-radiated back to obtain the observed temperature. To be clear,  $f$  is just a  
 122 fractional parameter. In 1950 it is some function of the GHGs (with no feedbacks). In 2019 it is more complex and  
 123 includes feedback effects. However, it is primarily related to GHGs re-radiation since  $P_{GHG} \approx P_{GHG'+Feedback}$ .

124  
 125 We then write

$$126 \quad P_{Total} = \sigma T^4 \text{ and } P_\alpha = \sigma T_\alpha^4 \quad (17)$$

127 We will find  $T_\alpha/T \approx \beta$ .

## 128 129 2.5 Balancing $P_{out}$ and $P_{in}$

131 Although Eq. 15 is reasonably simple, it turns out that  $f_1$  has a uniquely defined value obtained when balancing the  
 132 energy.

### 134 2.5.1 Balancing $P_{out}$ and $P_{in}$ in 1950

136 In order to balance the energy in with the energy out in 1950 with no global warming imbalance we can still start  
 137 with Eq. 15. In equilibrium the radiation that leaves must balance what comes in  $P_\alpha$  so that

$$139 \quad \begin{aligned} Energy_{out} &= (1-f_1)P_\alpha + (1-f_1)P = (1-f_1)P_\alpha + (1-f_1)\{P_\alpha + f_1P_\alpha\} \\ &= (1-f_1)\{2P_\alpha + f_1P_\alpha\} = 2P_\alpha - f_1P_\alpha - f_1^2P_\alpha = Energy_{in} = P_\alpha \end{aligned} \quad (18)$$

140  
 141 In 1950 the value of  $f$  solves the quadratic equation

$$143 \quad f_1^2 + f_1 - 1 = 0 \quad (19)$$

144 This yields the unique value in 1950

$$145 \quad f_1 = 0.618 \quad (20)$$

### 146 2.5.2 Warming Imbalance in 2019

148 The re-radiation parameters  $f_1$  and  $f_2$  are connected and from Eq. 15 and 16 we have

$$150 \quad f_2 = f_1 + \left( \frac{P_{2019}}{P_\alpha} - \frac{P_{1950}}{P_\alpha} \right) = f_1 + \Delta f \quad (21)$$

151 In this way  $f_2$  is a function of  $f_1=0.618$  and the differences in the global warming residuals can be defined as  $\Delta f$ .

## 153 3.0 Results and Discussion

155 Since the re-radiation parameter  $f_1=0.618$ , in order to obtain  $T_{1950}=13.89^\circ\text{C}$  ( $287.038^\circ\text{K}$ ), the only adjustable  
 156 parameter in our simple model is the Earth's albedo. This value requires an albedo value of 0.3008 (see Table 1) to  
 157 obtain the correct value  $T_{1950}$ . This is a reasonable and similar to values cited in the literature [11].

159 In 2019, the average temperature of the Earth is  $T_{2019}=14.84^\circ\text{C}$  ( $287.99^\circ\text{K}$ ). Here we are not sure of the albedo since  
 160 it likely changed due to UHI increase, snow and ice melting and cloud coverage changes. The IPCC value in AR5  
 161 [6] is 0.294118. However, this would represent a 3% change since 1950 which may be an overestimation. In our  
 162 assessment, we will assume a 1% change. Then the  $f_2$  parameter is adjusted to 0.6324 in order to obtain  $T_{2019}$ .  
 163 Results are provided in the Table 1. The results yields  $P_{Total\_1950}=384.918 \text{ W/m}^2$  and  $P_{Total\_2019}=390.024 \text{ W/m}^2$ . We  
 164 find that

$$165 \quad \Delta P_{Total} = P_{2019} - P_{1950} = 5.097 \text{ W/m}^2 \quad (22)$$

166 and

$$167 \quad \Delta T_{Total} = T_{2019} - T_{1950} = 0.95^\circ\text{C} \quad (23)$$

168 which is the observed surface temperature increase since 1950.

169  
170 **Table 1** Model results

Year	T( $^\circ\text{K}$ )	$T_\alpha$ ( $^\circ\text{K}$ )	$f_1, f_2$	$\alpha, \alpha'$	$P_\alpha, P_{\alpha'}$ ( $\text{W/m}^2$ )	$P_{GHG}$ ( $\text{W/m}^2$ ) $P_{GHG'+feedback}$	$P_{Total}$ ( $\text{W/m}^2$ )
2019	287.989	254.78	0.6324	29.779	238.927	149.870	390.024
1950	287.0395	254.51	0.618	30.08	237.903	147.024	384.918
$\Delta 2020-1950$	<b>0.95</b>	0.27	<b>1.44%</b>	-0.3 (1%)	<b>1.024</b>	2.846	<b>5.097</b>

171  
172 The table below summarizes model results for the specified albedos and setting the model to the observed Earth's  
173 surface temperatures.

174  
175 To show model consistency, the forcing change  $5.097 \text{ W/m}^2$  resulting in a  $0.95^\circ\text{K}$  rise, should agree with what is  
176 expected from Planck's feedback parameter. From Eq. 14 it is evident that

$$177 \beta^4 \Delta R_{\text{TOA}} = 5.097 \times \beta^4 = 3.132 \text{ W/m}^2 \quad (24)$$

178  
179 This illustrates the consistency of the simple re-radiation model. Then Planck's feedback parameter temperature rise  
180 is in agreement with what is observed

$$181 3.139 \text{ W/m}^2 \times (1/3.3)^\circ\text{K/W/m}^2 = 0.949^\circ\text{K at } T_s \quad (25)$$

### 182 3.1 Why the Re-radiation Parameter is Significant

183  
184 In Table 1, the measure of  $\Delta f = 1.44\%$  fractional increase is due to re-radiation change. This is significant. From Eq.  
185 15, 16 and 21 we can illustrate this key characteristic of the climate change

$$186 \Delta f = \left( \frac{P_{2019}}{P_{\alpha'}} - \frac{P_{1950}}{P_{\alpha}} \right) = \left( \frac{P_{\text{GHG}'+F}}{P_{\alpha'}} - \frac{P_{\text{GHG}}}{P_{\alpha}} \right) \approx \left( \frac{P_{\text{GHG}'+F} - P_{\text{GHG}}}{P_{\alpha}} \right) \quad (26)$$

187  
188 Therefore  $f$  is an estimate of climate re-radiation and  $\Delta f$  an estimate of climate change from a different perspective.  
189 It is a measure of GHG increase and the feedback relative to the initial radiation, and is generally helpful in looking  
190 at how our climate is working. Furthermore, we can deduce an albedo advantage.

### 191 3.2 The Albedo Advantage

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193 We can look at an important ratio, the power created by the albedo effect compared to GHGs in 1950. The initial  
194 radiation is  $P_{\alpha}$  which heats the Earth to  $254.51\text{K}$  then according to Eq. 15 and Table 1, the  $P_{\text{GHG}}$  energy originates  
195 from a fraction of this original heating due to re-radiation as  $fP_{\alpha}$

$$196 \frac{P_{\alpha} + P_{\text{GHG}}}{P_{\text{GHG}}} = \frac{P_{\alpha} + fP_{\alpha}}{P_{\text{GHG}}} = \frac{P_{\alpha} + fP_{\alpha}}{fP_{\alpha}} = \frac{1 + f_1}{f_1} = \frac{1.62}{0.62} = 2.62 \quad (27)$$

197  
198 In general, this also means that albedo change has a higher impact factor in climate forcing, 2.6 times larger than  
199  $\Delta P_{\text{GHG}}$  as well, that is a change,  $\Delta P_{\alpha}$  compared with a change in  $\Delta P_{\text{GHG}}$  would yield the same impact factor

$$200 \frac{\Delta P_{\alpha} + \Delta P_{\text{GHG}}}{\Delta P_{\text{GHG}}} \approx \frac{\Delta P_{\alpha} + f \Delta P_{\alpha}}{f \Delta P_{\alpha}} \approx \frac{1 + f_1}{f_1} = \frac{1.62}{0.62} = 2.62 \quad (28)$$

201  
202 This is a key reason that UHIs, cloud coverage, snow and ice melting, can create significant climate effects.  
203 Appendix A puts this important impact factor in layman's terms.

204  
205 In this view, an albedo solution is advantageous having significant potential for reversing global warming or  
206 ignoring it, as in UHIs likely can create serious issues. Therefore, trying to control global warming by reducing  
207 GHGs is important. However, certainly an albedo approach is more advantageous. It reduces both initial absorption  
208 and its potential for its re-radiation. Its impact rating can be taken as 162% compared to re-radiation  $f$  with a 62%  
209 impact by comparison according to Eq. 27 and 28, yielding a 2.6 times higher advantage. It is important to realize  
210 that because the albedo solution can highly impact GW and reverse trends, it is also vital in preventing a tipping  
211 point from occurring.

### 212 3.3 Planck-Albedo Feedback Parameter

213  
214 The albedo changes in Table 1, is:  $\% \Delta \alpha = 1\%$ . The albedo  $\Delta P_{\alpha}$  change in Table 1 is  $1.024 \text{ W/m}^2$ . We note that we  
215 can define a unique Planck-Albedo parameter  $\lambda_{\% \Delta \alpha} = \Delta P_{\alpha} / \% \Delta \text{albedo}$ . To illustrate from Table 1

$$216 \lambda_{\% \Delta \alpha} = 1.024 \text{ W/m}^2 / 1\% = 1.024 / 1\% \quad (29)$$

217  
218 This parameter can also be expressed per degree (noting the  $0.95^\circ\text{K}$  change in Table 1)

$$\lambda_{\% \Delta \alpha \Delta T} \approx 1W / m^2 / \Delta \% albedo / ^\circ K \quad (30)$$

229  
230 The parameter was first noted in Feinberg 2020 [2] but is featured here as a modeling tool. We term it the Planck-  
231 Albedo parameter, since it relates to blackbody ( $P_\alpha$ ) absorption. A simple numeric example is given in the  
232 Conclusion to illustrate how it provides helpful estimates. This interesting parameter arises from the basic  
233 assessment

$$\lambda_{\% \Delta \alpha} = \frac{(\Delta E_o)_\alpha}{\alpha_1 - \alpha_2} \frac{100}{\alpha_1} = \frac{E_o (\alpha_1 - \alpha_2)}{\alpha_1 - \alpha_2} \frac{100}{\alpha_1} = E_o \alpha_1 / 100 \approx 1W / m^2 / \% \Delta albedo \quad (31)$$

235 where  $E_o = 340 \text{ W/m}^2$  and when  $\alpha_1$  is 29.4118%, the value  $1.000 \text{ W/m}^2 / \Delta \% albedo$  is obtained. We note the value  
236 29.4118% (100/340) is given in AR5 [6]. The parameter's relationship to  $\lambda_\alpha$  is

$$\lambda_\alpha = \lambda_{\% \Delta \alpha \Delta T} x \% \Delta \alpha \quad (32)$$

239 and the feedback parameter including  $f$  re-radiation is in 2019

$$\lambda_\alpha^\dagger = \lambda_{\% \Delta \alpha \Delta T} x \% \Delta \alpha x 1.618 \quad (33)$$

#### 244 4.0 Conclusion

246 In this paper we provided a simple re-radiation global warming model. The model shows consistency with the  
247 Planck parameter. We noted that the re-radiation parameter increased by about 1.44% due to global warming from  
248 1950 to 2019, illustrating the warming from a different perspective. From the model, the albedo effect was  
249 quantified having an impact rating of 162% compared to GHGs with 62%. The albedo effect then yields a 2.6 times  
250 higher advantage upon comparison. These results strongly support moving forward with solar geoengineering  
251 solutions [2, 7-9].

253 We also found a handy parameter that we termed the Planck-Albedo parameter which is about  
254  $\lambda_{\% \Delta \alpha \Delta T} \approx 1W / m^2 / \Delta \% albedo / ^\circ K$ . This can be helpful in quickly estimating the effect of an albedo change on global  
255 warming and in assessing  $\lambda_\alpha$ . For example, Feinberg 2020 [1] suggested a goal of 1.5% geoengineering albedo  
256 change. Using this parameter, an impact of 1.5 Watts/m<sup>2</sup> warming reduction should result. Given a 1.62 reemission  
257 factor (Eq. 28), this is 2.4W/m<sup>2</sup> improvement. With a reduction in water-vapor feedback, often estimated by a factor  
258 of 2 [10], provides an overall resulting effect that could be as high as 4.8W/m<sup>2</sup>. Feasibility is discussed in more  
259 detail in Feinberg's 2020 paper [1] and other solutions have been proposed [6-9].

#### 261 Appendix A: Quantifying the Albedo Advantage in Layman's Terms

263 It may be helpful for the reader to have a layman's view of how the 2.62 factor comes about. Consider the Earth  
264 with a roof. The roof represents the GHGs over the Earth and only allows 40% of any energy leaves with the rest  
265 returning to Earth. Sunlight comes in and some is absorbed and heats the Earth's floor to 255°K (-2.3°F very cold).  
266 Let's say it takes 100 units of energy. The heat rises but only 40 units of energy can leave so 60 units comes back  
267 and warms the Earth's floor some more to 288°K (57°F average temp of Earth). On average the Earth's floor is  
268 warmed a total of 160 units. The Sun keeps warming the Earth's floor at 100 units on average and the roof keeps  
269 sending back 60. So the roof is responsible for 60 units on average of energy and the Earth's floor is warmed up to  
270 160 units on average. We can write this as

$$272 \text{ Energy units: } 160 = 100 + 60 = 100 + 100 \times 0.6$$

274 We see the 100 units is in two places in the equation, while the 60 is only in one place. That is without the floor  
275 absorption first, the roof cannot keep the Earth warm. Therefore, the heat coming from the Earth's floor results  
276 in 160 units and the roof is only 60 units by comparison. The impact factor is

- 278 • 160/60=2.66, that is the heat from the Earth's floor has this much larger impact.

280 Alternately, for every unit of energy given off, by the Earth's floor after absorption it is equivalent to causing 1.6  
281 units of heating while the roof (GHG) is only responsible for 0.6.

283 How much heat leaves in equilibrium? There was the initial 40 leaving of the 100 units of energy absorbed and  
284 radiated. As well the Earth's floor received a total of 160 units but the roof only let 40% leave that is another 64  
285 (=0.4 x 160) units of energy leaving. The total leaving is 104 units in equilibrium so roughly 100 units comes in and  
286 almost same goes out.

287  
 288 This can be refined to 61.8% (Eq. 20). Then 100 units is absorbed and radiated, then 38.2 units initially leave, and  
 289 61.8 units is radiated so the Earth's floor is heated to 161.8 units of energy. From this  $0.382 \times 161.8$  leaves=61.8  
 290 units or energy. The total is  $61.8+38.2=100$  units of energy leaves and another 100 units comes and equilibrium is  
 291 established. Any difference causes global warming.

- 292  
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